

# Simulation and observation for volume emission rates emitted from O<sub>2</sub>(0-1) and O(<sup>1</sup>S) nightglow in Northwest China

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Received 6 November 2018; revised 18 December 2018; accepted 19 December 2018; posted 20 December 2018 (Doc. ID 351182); published 0 MONTH 0000

Being susceptible to the change of atmospheric conditions, the volume emission rate (VER) is very suitable to be used as a light source by passive remote sensing for measuring atmospheric wind and temperature. Thus, the VERs emitted from O<sub>2</sub>(0-1) and O(<sup>1</sup>S) of the nightglow at 80–120 km are studied in this paper. Based on the Naval Research Laboratory Mass Spectrometer Incoherent Scatter (NRLMSISE-00) model data and the ground-based airglow imaging interferometer (GBAII) instrument observation for a local time and place, simulated VER profiles represented by four layers are obtained for the nightglow of O<sub>2</sub>(0-1) and O(<sup>1</sup>S). The O<sub>2</sub>(0-1) nightglow model peak values at 94 km on 6 December 6 2013 and 8 November 2011 are 8111 photons · cm<sup>-3</sup> · s<sup>-1</sup> and 8406 photons · cm<sup>-3</sup> · s<sup>-1</sup>, respectively; however, the O(<sup>1</sup>S) VER peak at a higher altitude of about 96 km on 18 December 2011 is only 338 photons · cm<sup>-3</sup> · s<sup>-1</sup>. The upper atmospheric VER values have been derived to transfer into the ground-based detected column intensities by our GBAII prototype. The calculated column integrated emission rates (IERS) of O<sub>2</sub>(0-1) for 0° and 45° zenith angles are 1.48 × 10<sup>7</sup> and 1.91 × 10<sup>7</sup> photons · cm<sup>-2</sup> · s<sup>-1</sup>, respectively; the calculated column IERS of O(<sup>1</sup>S) are 5.53 × 10<sup>5</sup> and 7.03 × 10<sup>5</sup> photons · cm<sup>-2</sup> · s<sup>-1</sup>, respectively. Correspondingly, the detected column IERS obtained by GBAII are 1.87 × 10<sup>7</sup> for O<sub>2</sub>(0-1) and 6.57 × 10<sup>5</sup> photons · cm<sup>-2</sup> · s<sup>-1</sup> for O(<sup>1</sup>S). © 2019 Optical Society of America

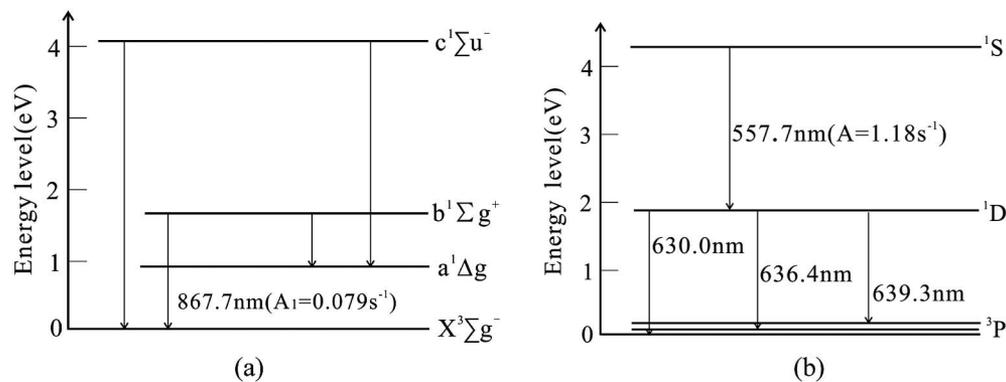
<https://doi.org/10.1364/AO.99.099999>

## 1. INTRODUCTION

Airglow is a low-light-level phenomenon caused by the solar electromagnetic radiation in the upper atmosphere of the Earth. When the direct or indirect solar radiation excites some atmospheric species to their higher energy level states, photons can be emitted by transitions from the excited states to the lower energy states due to instability. Because there is a longer lifetime at the excited states and the moving atmosphere species, the photons can run up to the balanced atmospheric emission, termed airglow, including dayglow and nightglow. In fact, as the airglow exists permanently in the upper atmosphere, it is very suitable for it to be used as a light source by passive remote sensing for measuring atmospheric wind and temperature. Although airglow is very difficult to detect due to its low light level intensity [1], the radiosonde, rocket, satellite measurement, and ground-based technologies are developed to detect airglow [2].

With these advanced instruments, such important parameters in the upper atmosphere as wind vector, temperature, and volume emission rate (VER) can be obtained. Also, these excellent spatiotemporal resolution data sets provide support for research on gravity waves, ozone concentration [3], and other atmospheric compositions [4].

The theory and technology for detecting atmospheric temperature by using airglow emission were first proposed by the York University of Canada group in 1966 [5]. The Wind Imaging Interferometer (WINDII) [6], a wide-angle Michelson interferometer (MI) on NASA's upper atmospheric research satellite (UARS, 1991–2005), was a representative instrument that was able to measure the upper atmospheric (80–300 km) wind vector, temperature, and VER through multiple airglow emission lines. The high-resolution Doppler imager (HRDI) [7], another instrument on UARS, was also used to detect



F1:1 **Fig. 1.** Energy level scheme of  $O_2(0-1)$  and  $O(^1S)$  nightglows.

61 the low-middle atmospheric (10–120 km) wind field  
 62 through Fabry–Perot interferometer (FPI) technology. The  
 63 observation regions of WINDII and HRDI had a 40 km  
 64 overlap in altitude, with the long-term observed results  
 65 from them showing excellent agreement by two different ap-  
 66 plied technologies (MI and FPI) [8]. The VER of  $O_2(0-0)$   
 67 airglow at 94 km altitude observed by WINDII was about  
 68  $4000\text{--}5000 \text{ photon} \cdot \text{cm}^{-3} \cdot \text{s}^{-1}$  [2], while the VER peak of  
 69  $O(^1S)$  nightglow was  $60\text{--}200 \text{ photon} \cdot \text{cm}^{-3} \cdot \text{s}^{-1}$  at 90–97 km  
 70 altitude (versus March 1993) [9]. The VER of  $O_2(0-0)$  airglow  
 71 at 94 km altitude was also obtained by HRDI [10]. The meso-  
 72 pause oxygen rotational temperature imager (MORTI), a  
 73 ground-based instrument made to study the airglow, had been  
 74 upgraded to measure VER  $O_2(0-1)$  and the vertically averaged  
 75 temperature of the  $O_2(0-1)$  atmospheric band nightglow layer  
 76 at 94 km altitude and OH (6-2) Meinel layer at 87 km [4]. The  
 77 Spectral Airglow Temperature Imager (SATI), an improved  
 78 version based on MORTI, was aimed at detecting rotational  
 79 temperature and VER [11]. Sounding the atmosphere using  
 80 broadband emission radiometry (SABER), another space-based  
 81 instrument on the thermosphere ionosphere mesosphere  
 82 energetics and dynamics (TIMED) satellite, obtained results  
 83 from OH (6-2) airglow and nitric oxide VER (100–180 km)  
 84 at  $5.3 \mu\text{m}$  wavelength [12].

85 The airglow intensity as well as its VER has strong latitu-  
 86 dinal, temporal, and local dependencies, and airglow detectors  
 87 need to be distributed in different representative areas. Some  
 88 local observations in northwest China ( $N34^\circ$ ,  $E108^\circ$ ) are pre-  
 89 sented in this paper. The ground-based airglow imaging inter-  
 90 ferometer (GBAII), which was proposed in earlier research by  
 91 the authors, was capable of simultaneously measuring the wind  
 92 vector and temperature at 90–100 km altitude [13,14]. GBAII,  
 93 an integrated design of a wide-angle MI and a F-P filter, can  
 94 employ both the “four intensity algorithms” [6] and “rotational  
 95 line measurement temperature” methods [3]. A liquid-crystal  
 96 on silicon (LCoS) is used as a mirror of the large air gap  
 97 wide-angle MI [15] to be the static stepped device [16]. The  
 98 nightglow sources of the  $O_2$  atmospheric (0-1) band centered at  
 99 867.7 nm (867.8, 867.7, 867.3, 867.2, 866.2, 866.8, 866.6,  
 100 866.3, 866.2, 865.8, 865.7, 865.4, and 865.2 nm, 12 lines)  
 101 and  $O(^1S)$  557.7 nm (single line) are selected to detect the  
 102 atmospheric wind vector and temperature by GBAII. The

103 VERs of  $O_2(0-1)$  and  $O(^1S)$  nightglow are presented in  
 104 subsequent sections of this paper.

105 The present paper is arranged as follows: Section 2 gives the  
 106 simulation results on the VER of  $O_2(0-1)$  and  $O(^1S)$  night-  
 107 glow based on the Barth mechanism; Section 3 describes the  
 108 calculated nightglow integrated emission rate (IER) results  
 109 for GBAII’s geometry; Section 4 details the field experiments  
 110 and the inversed nightglow IER results by using GBAII;  
 111 Section 5 presents some discussions as well as a concluding  
 112 summary.

## 113 2. SIMULATED NIGHTGLOW VER OF $O_2(0-1)$ AND $O(^1S)$ 114

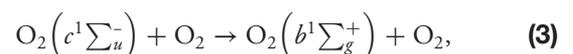
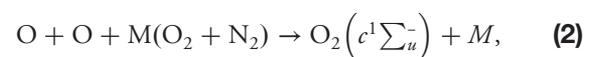
115 For a certain atmospheric species, VER refers to the number  
 116 of emitted photons per second and volume from its higher  
 117 energy level states to lower energy states. VER’s unit is  
 118  $\text{photons} \cdot \text{cm}^{-3} \cdot \text{s}^{-1}$ , and its expression is

$$119 \text{VER} = A \cdot [c], \quad (1)$$

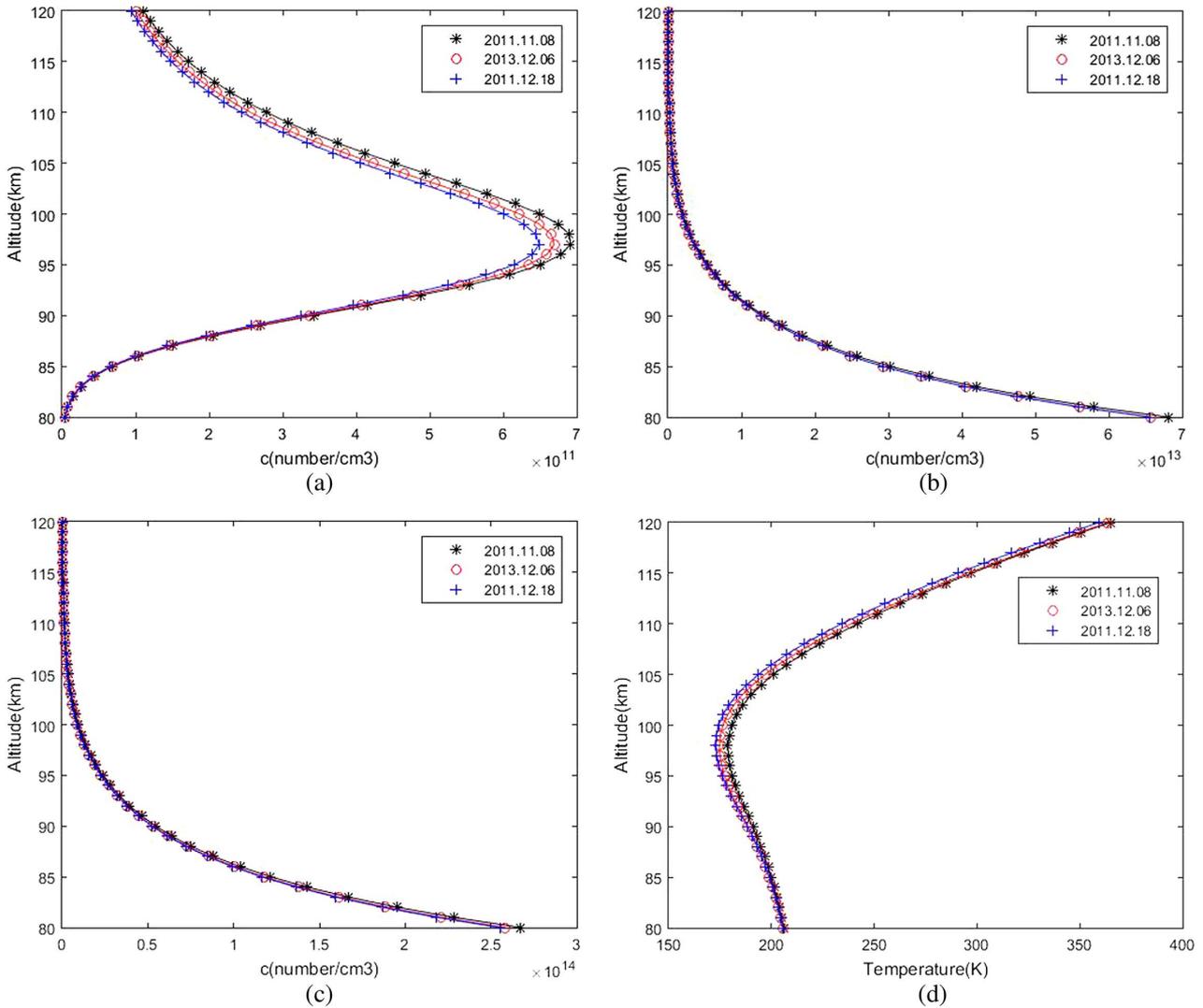
120 where  $A$  denotes spontaneous radiation Einstein-A coefficient  
 121 and  $[c]$  denotes the atmospheric species concentration (number  
 122 densities) like  $[O_2(0-1)]$  or  $[O(^1S)]$ .

### 122 A. VER of $O_2(0-1)$ Nightglow

123 Molecular oxygen  $O_2(0-1)$  867.7 nm nightglow is the spec-  
 124 trum transfer of ( $b^1\Sigma_g^+ \rightarrow X^3\Sigma_g^-$ ) shown in Fig. 1(a), where  
 125  $X$  is the ground state,  $b$  is the second excited state,  $\Sigma$  is the  
 126 molecular excited state,  $g$  is the vibronic level symmetry, and  
 127  $u$  is the vibronic level antisymmetry. According to the Barth  
 128 mechanism, the photochemical emission reaction involved  
 129 in the  $O_2$  atmospheric band spectrum can be described in the  
 130 following two-step process [2]:



133 where the quenching of  $O_2(b^1\Sigma_g^+)$  by the atmospheric  $O_2$   
 and  $N_2$  is included, and  $M$  is set to be equal to the molecules  
 of  $[O_2]$  and  $[N_2]$ . For the two-step transfer of the Barth



**Fig. 2.** Vertical concentration profiles mapping–altitude: (a) atoms oxygen concentration, (b) molecule O<sub>2</sub> concentration, (c) molecule N<sub>2</sub> concentration, (d) upper atmospheric temperature.

mechanism, the VER for this process can be defined and evaluated by the following predecessor [17]:

$$\begin{aligned}
 & \text{VER}_{\text{O}_2(0-1)} \\
 &= \frac{k_r A_1 [\text{O}]^2 ([\text{O}_2] + [\text{N}_2]) [\text{O}_2]}{(A_2 + K_2^{\text{O}_2} \cdot [\text{O}_2] + K_2^{\text{N}_2} \cdot [\text{N}_2]) \cdot (7.5[\text{O}_2] + 33[\text{O}])}, \tag{4}
 \end{aligned}$$

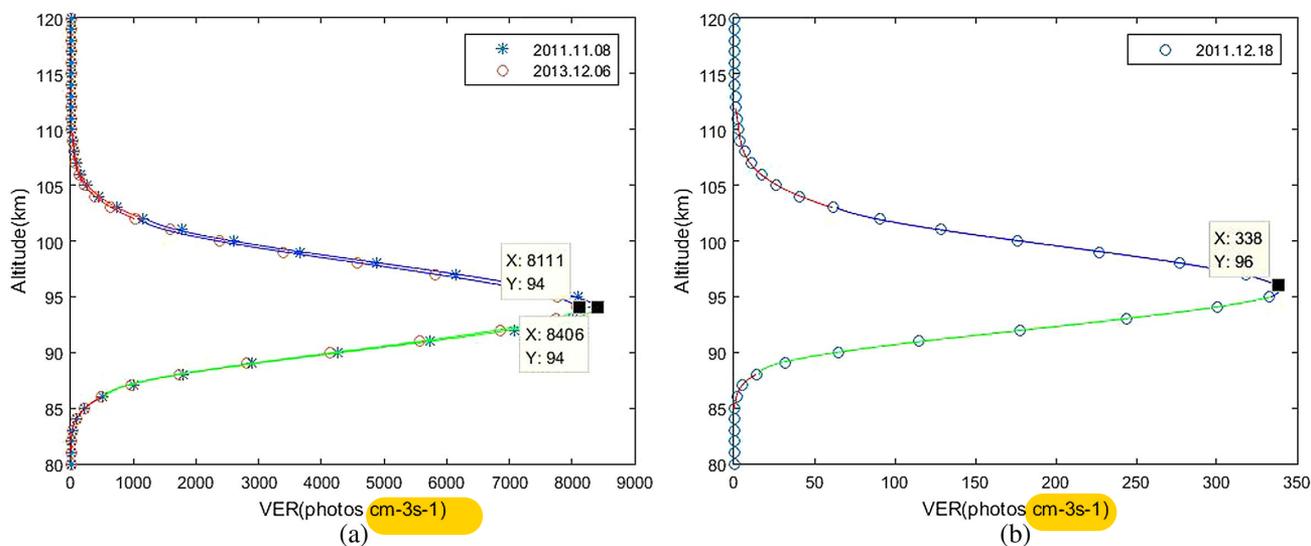
where  $k_r$  is the rate coefficient for the three-body recombination of the O atom;  $k_r = 4.7 \times 10^{-33} (300/T)^2 \text{ cm}^6 \cdot \text{s}^{-1}$ .  $K_2^{\text{O}_2} = 4 \times 10^{-17} \text{ cm}^3 \cdot \text{s}^{-1}$  and  $K_2^{\text{N}_2} = 2.2 \times 10^{-15} \text{ cm}^3 \cdot \text{s}^{-1}$  are the rate coefficients for quenching by O<sub>2</sub> and N<sub>2</sub>, respectively;  $A_1$  and  $A_2$  are the transition probabilities of (0-0) band and O<sub>2</sub>(b<sup>1</sup>Σ<sub>g</sub><sup>+</sup>) ( $A_1 = 0.079 \text{ s}^{-1}$  and  $A_2 = 0.083 \text{ s}^{-1}$ ), respectively; [O], [O<sub>2</sub>], and [N<sub>2</sub>] are, respectively, the number densities of atomic oxygen, molecular oxygen, and nitrogen.

The data of concentrations of atomic [O], molecular [O<sub>2</sub>], and [N<sub>2</sub>] as well as the upper atmospheric temperature profile

of the Naval Research Laboratory Mass Spectrometer Incoherent Scatter (NRLMSISE-00) of the Community Coordinated Modeling Center (CCMC) are used in this paper [18]. According to the GBAll's parameters and the local observation time and place for the nightglows, for example, the experimental date (8 November 2011) and the geographical position of Xi'an (N34°13'23", E 108°59'39") in China, the discrete concentration or number densities can be obtained and are shown in Fig. 2. After bringing these number densities of [O], [O<sub>2</sub>], [N<sub>2</sub>] and corresponding atmospheric temperature into Eq. (4), a group of discrete VER values can be obtained, which are shown in Fig. 3(a). The VER peak values of VER were about 8111 and 8406 photons · cm<sup>-3</sup> · s<sup>-1</sup> on 6 December 2013 and 8 November 2011, respectively, at 94 km altitude.

Because the VER profiles mapped to the altitude shown in Fig. 3 are Gaussian profiles, it is troublesome for the ground detected instrument, so we give a so-called "four-layer method" to fit the Gaussian profile. The aim is to transform the integral

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F3:1 **Fig. 3.** Nightglow VER profile mapping–altitude: (a) O<sub>2</sub>(0–1) 867.7 nm nightglow, (b) O(<sup>1</sup>S) 557.7 nm nightglow.

**Table 1.** Fitting Coefficients versus Eq. (5) of O<sub>2</sub>(0–1) VER Profile for Two Dates

Altitude	8 November 2011				6 December 2013			
	$a_1$	$a_2$	$a_3$	$a_4$	$a_1$	$a_2$	$a_3$	$a_4$
81 < $h$ < 86 km	5.4145	$-1.3236 \times 10^3$	$1.0785 \times 10^5$	$-2.9292 \times 10^6$	5.2482	$-1.2830 \times 10^3$	$1.0454 \times 10^5$	$-2.8394 \times 10^6$
86 < $h$ < 94 km	-28.219	$7.6283 \times 10^3$	$-6.8593 \times 10^5$	$2.0521 \times 10^6$	-27.786	$7.5089 \times 10^3$	$-6.7502 \times 10^5$	$2.0190 \times 10^7$
94 < $h$ < 102 km	22.410	$-6.5978 \times 10^3$	$6.4622 \times 10^5$	$-2.1052 \times 10^7$	21.747e	$-6.3960 \times 10^3$	$6.2583 \times 10^5$	$-2.0367 \times 10^7$
102 < $h$ < 110 km	-3.6638	$1.1919 \times 10^3$	$-1.2928 \times 10^5$	$4.6743 \times 10^6$	-3.4401	$1.1182 \times 10^3$	$-1.2117 \times 10^5$	$4.3773 \times 10^6$

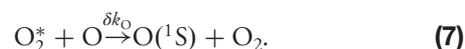
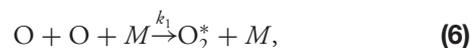
165 VER value conveniently along the line of sight for the column  
 166 IER detected by the ground-based CCD detector. In later  
 167 Section 3.B, a local continuous series of powers of the VER  
 168 profile needs to be fitted with the results shown in Fig. 3(a).  
 169 Of course, at the NRLMSISE window condition, our Xi'an  
 170 local time 2:00 am and location of longitude and latitude  
 171 (N34°, E108°) are selected as perfectly as possible to match  
 172 GBAIL's observation conditions. The VER profile is a function  
 173 of altitude  $h$  and is expressed as

$$\text{VER}(h) = a_1 h^3 + a_2 h^2 + a_3 h + a_4, \quad (5)$$

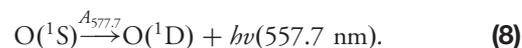
174 where  $a_1$ ,  $a_2$ ,  $a_3$ , and  $a_4$  are different coefficients. For the best  
 175 fitting degree,  $h$  is divided into four regions, i.e., 81–86 km,  
 176 86–94 km, 94–102 km, and 102–110 km O<sub>2</sub>(0–1) nightglows  
 177 shown in Fig. 3(a). The corresponding coefficients of  $a_1$ ,  $a_2$ ,  $a_3$ ,  
 178 and  $a_4$  for two days are shown in Table 1. The fitting degree is  
 179 99.0% according to the least-squares method.

### 180 B. VER of O(<sup>1</sup>S) Nightglow

181 According to the Barth mechanisms, the atomic oxygen  
 182 O(<sup>1</sup>S)557.7 nm nightglow emission transition shown in  
 183 Fig. 1(b) can be expressed as the transition of O(<sup>1</sup>S) → O(<sup>1</sup>D).  
 184 In order to calculate the VER of O(<sup>1</sup>S), the O(<sup>1</sup>S) concentration  
 185 must be known for Eq. (1). By following three photochemical  
 186 reaction actions, the concentration can be obtained as follows:



For a balanced atmosphere, generation rate is equal to the  
 consumption rate, and thus there is the following equilibrium  
 equation:



The VER of O(<sup>1</sup>S) can be written as [18] 190

$$\text{VER}_{\text{O}(\text{1S})} = \frac{A_{557.7 \text{ nm}} k_1 [\text{O}]^3 [M]}{(A(\text{1S}) + k_2 [\text{O}_2]) (C_0 + C_1 [\text{O}] + C_2 [\text{O}_2])}, \quad (9)$$

where  $C_0 = A(\text{O}_2^*)/\beta\delta k_{\text{O}}$ ,  $C_1 = 1/\beta\delta$ ,  $C_2 = (k_{\text{O}_2} + Rk_{\text{N}_2})/\beta\delta k_{\text{O}}$ ,  
 $R = [\text{N}_2]/[\text{O}_2]$ , and  $M = [\text{O}_2] + [\text{N}_2]$ . The values of the parameters are  
 $k_1 = 4.7 \times 10^{-33} (300/T)^2 \text{ cm}^6 \text{ s}^{-1}$ ,  
 $k_2 = 4.0 \times 10^{-12} \exp(-867/T) \text{ cm}^3 \text{ s}^{-1}$ ,  
 $A_{557.7 \text{ nm}} = 1.18 \text{ s}^{-1}$ ,  
 $A(\text{1S}) = 1.35 \text{ s}^{-1}$ ,  $C_0 = 0$ ,  $C_1 = 211$ , and  $C_2 = 15$  [18], respectively.  
 Similarly, the discrete and fitted VER values of O(<sup>1</sup>S) 557.7 nm  
 nightglow can be obtained [see Fig. 3(b)] after substituting the  
 number densities of atomic [O] and molecular [O<sub>2</sub>] into Eq. (10)  
 from NRLMSISE-00 [19]. The peak value of the VER was about  
 338.3 photons · cm<sup>-3</sup> · s<sup>-1</sup> on 18 December 2011 at 96 km  
 altitude. As O<sub>2</sub>(0–1) nightglow, we define four

**Table 2. Nightglow’s Transmittance versus the Wavelength and the Zenith Angle**

	Wavelength (nm)	Transmittance (0° zenith angle)	Transmittance (45° zenith angle)	Selected Parameters from MODTRAN Model	Date
T2:1					
T2:2	867.7	0.670	0.567	Urban visibility of 10 km, ground temperature of 278 K	8 Nov. 2011
T2:3	867.7	0.671	0.560	Urban visibility of 10 km, ground temperature of 270 K	6 Dec. 2013
T2:4	867.7	0.833	0.758	Rural visibility of 23 km, ground temperature of 278 K	18 Apr. 2018
T2:5	557.7	0.658	0.551	Rural visibility of 23 km, ground temperature of 273 K	18 Dec. 2011

202 homologous altitude  $h$  regions, and the corresponding VER  
 203 expressions are directly obtained by MATLAB as follows. (The  
 204 recent corresponding MODTRAN data is only updated by the  
 205 deadline on 17 April 2017, so we only give one simulation equa-  
 206 tion). The fitting degree is 98.8% according to the least-squares  
 207 method. The VER expressions are

$$\begin{cases} \text{VER}_1(h) = 0.5023h^3 - 1.2845 \times 10^2 h^2 + 1.0950 \times 10^4 h - 3.1118 \times 10^5 & (85 \leq h < 88 \text{ km}), \\ \text{VER}_2(h) = -1.3727h^3 + 3.7895 \times 10^2 h^2 - 3.4809 \times 10^4 h + 1.0640 \times 10^6 & (88 \leq h < 95 \text{ km}), \\ \text{VER}_3(h) = 1.0916h^3 - 3.2613 \times 10^2 h^2 + 3.2426 \times 10^4 h - 1.0727 \times 10^6 & (95 \leq h < 103 \text{ km}), \\ \text{VER}_4(h) = -0.1410h^3 + 4.6622 \times 10^1 h^2 - 5.1381 \times 10^3 h + 1.8878 \times 10^5 & (103 \leq h \leq 112 \text{ km}), \end{cases} \quad (10)$$

gathered together by the GBAlI detector, which has a field of  
 view (FOV) incident angle of 6°, and the slanting optical path  
 nightglow for the 45° zenith angle of incident light [13,14] is  
 shown as the green region in Fig. 4. Thus, the nightglow IER  
 detected by GBAlI’s CCD detector can receive the nightglow  
 photons from the ground to 120 km altitude as follows:

208 **3. CALCULATED NIGHTGLOW IER FOR GBAlI**

210 **A. Nightglow Transmittance**

211 When the nightglow emission at 80–120 km altitude arrives at  
 212 our ground-based GBAlI detector after penetrating the atmos-  
 213 pheric layer below, its IER will be reduced due to the atmos-  
 214 pheric scattering and absorption. The atmospheric molecules’  
 215 absorption and scattering at 867.7 nm and 557.7 nm are due  
 216 to water vapor, carbon dioxide, and aerosols. In fact, the  
 217 molecules’ absorption and scattering can be ignored [20], and  
 218 the atmospheric attenuation factor is considered more in the  
 219 atmosphere below 10 km altitude, so the scattering may be con-  
 220 sidered to be affected by the aerosol influence for the dusty  
 221 climates of Northwest China where GBAlI’s experiment is  
 222 performed.

223 Although the wavelength 867.7 nm is just at the atmos-  
 224 pheric window, the atmospheric transmittance must be  
 225 calculated when the nightglow enters the GBAlI’s detector.  
 226 The MODTRAN (moderate resolution atmospheric transmis-  
 227 sion) [21] is used as a reference for the nightglow transmittance.  
 228 Appropriate parameters for the GBAlI are selected as input for  
 229 MODTRAN. We set the “Atmosphere Model” parameter to be  
 230 mid-latitude summer and winter, the “Ground Temperature”  
 231 parameter to be the actual experiment temperature, and the  
 232 “Aerosol Models” parameter to be rural and urban. According  
 233 to the GBAlI’s different observation place, the “Visibility”  
 234 parameter was 23 km and 10 km, and the “Sensor Zenith”  
 235 was 180° and 135°. The nightglow transmittance results are  
 236 listed in Table 2.

237 **B. Calculation Nightglow IER for GBAlI**

238 As a ground-based station device, the GBAlI is able to accept  
 239 the full photon column of the nightglow intensity profile at  
 240 80–120 km as shown in Fig. 3. The nightglow emission is

$$\text{IER}_C = \int_{r_1}^{r_2} \int_{\theta=0^\circ}^{\theta=6^\circ} \int_{\varphi=0}^{2\pi} \frac{\text{VER}}{4\pi r^2} \tau \cdot r^2 \sin \theta dr d\theta d\varphi, \quad (11)$$

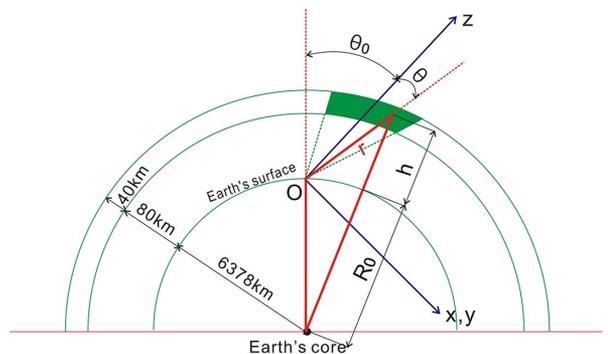
247 where  $I_{\text{IER}}$  is the calculated number of photons accepted by the  
 248 GBAlI’s CCD per area,  $\tau$  is the atmospheric transmittance,  
 249  $\text{VER}(h)$  is as mentioned earlier in Eq. (5),  $r$  is the slanting  
 250 optical path from the GBAlI to the nightglow species, and  
 251  $r_1 = 80 \text{ km}$ ,  $r_2 = 120 \text{ km}$ , and  $\text{VER}(h)$  are the different  
 252 expressions followed by Eq. (5) versus Table 1 and Eq. (11).  
 253 There is the following relation between  $r$  and  $h$  (see Fig. 4):

$$(R_0 + h)^2 = r^2 + R_0^2 - 2rR_0 \cos(\pi - \theta_0 - \theta), \quad (12)$$

254 where  $R$  denotes the Earth’s radius,  $h$  denotes altitude from 80  
 255 to 120 km,  $R_0 = R + h$ ,  $\theta_0$  denotes the GBAlI’s zenith angle  
 256 of incidence, and  $\theta$  denotes the GBAlI’s FOV. As  $R_0 \gg h$ ,  
 257 Eq. (12) can be simplified to

$$h = r \cos(\theta_0 + \theta). \quad (13)$$

258 According to Eqs. (5), (11), and (13), the  $\text{IER}_C$  can be



**Fig. 4.** Airglow contribution height for the GBAlI’s measurement geometry.

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F4:2

**Table 3. Nightglow IER Results for GBAIL<sup>a,b</sup>**

	Observation Date	Nightglow	Calculated IER <sub>C</sub> (0° zenith angle)	Calculated IER <sub>C</sub> (45° zenith angle)	Detected IER <sub>D</sub> (45° zenith angle)	Error
T3:1						
T3:2	8 Nov. 2011	O <sub>2</sub> (0–1)	1.48 × 10 <sup>7</sup>	1.91 × 10 <sup>7</sup>	1.87 × 10 <sup>7</sup>	2.1%
T3:3	6 Dec. 2013	O <sub>2</sub> (0–1)	1.42 × 10 <sup>7</sup>	1.79 × 10 <sup>7</sup>	1.27 × 10 <sup>7</sup>	29.0%
T3:4	18 Apr. 2018	O <sub>2</sub> (0–1)	No data	No data	1.29 × 10 <sup>7</sup>	
T3:5	18 Dec. 2011	O(1S)	5.53 × 10 <sup>5</sup>	7.03 × 10 <sup>5</sup>	6.57 × 10 <sup>5</sup>	6.5%
T3:6	18 Apr. 2018	O(1S)	No data	No data	1.347 × 10 <sup>6</sup>	

<sup>a</sup>Unit: photons · cm<sup>-2</sup> · s<sup>-1</sup>.

<sup>b</sup>Remarks: “No data” is because MODTRAN data is only updated to 17 April 2017.

$$\text{IER}_C = \int_{\theta=0}^{6^\circ} \int_{r_1}^{r_2} \int_{\varphi=0}^{2\pi} \frac{\text{VER}[r \cos(\theta_0 + \theta)] \cdot 10^6}{4\pi r^2} \cdot \tau \cdot r^2 \sin \theta d\varphi dr d\theta. \quad (14)$$

Bringing Eq. (10) into Eq. (14), the nightglow IER<sub>C</sub> is

$$\text{IER}_C = \frac{\tau}{2} \cdot 10^5 \cdot \left( \frac{a_1}{4} b^4 + \frac{a_2}{3} b^3 + \frac{a_3}{2} b^2 + a_4 b \right) \Bigg|_{80}^{120} \cdot \left[ -\cos \theta_0 \ln \frac{\cos(\theta_0 + 6)}{\cos \theta_0} - 6 \sin \theta_0 \right], \quad (15)$$

where the unit of IER<sub>C</sub> is photons · cm<sup>-2</sup> · s<sup>-1</sup>, the unit of *b* is denoted by kilometers (km), and the factor of 10<sup>5</sup> is the unit result from. Using *t*, *τ*, and *a*<sub>1</sub>–*a*<sub>4</sub> for each experiment into Eq. (16), the calculated IER of nightglow can be obtained. For example, for O<sub>2</sub>(0–1) 867.7 nm nightglow, let *τ* = 0.658 and 0.551 (see Table 2), and bringing the VER (*b*) functions in Eq. (5) versus Table 1 and Eq. (10) into Eq. (15), with the corresponding atmospheric transmittance listed in Table 2, the nightglow IER can be calculated. The results are 1.91 × 10<sup>7</sup> (photons · cm<sup>-2</sup> · s<sup>-1</sup>) and 1.48 × 10<sup>7</sup> (photons · cm<sup>-2</sup> · s<sup>-1</sup>) when *θ*<sub>0</sub> = 45° and *θ*<sub>0</sub> = 0°, respectively. For O(1S) 557.7 nm nightglow, when *τ* = 0.658 and 0.551 (see Table 2), its calculated intensity is 7.03 × 10<sup>5</sup> (photons · cm<sup>-2</sup> · s<sup>-1</sup>) for *θ*<sub>0</sub> = 45°, and the calculated IER is 5.53 × 10<sup>5</sup> (photons · cm<sup>-2</sup> · s<sup>-1</sup>) for *θ*<sub>0</sub> = 0°. [For the other data, see Columns 3 and 4 in Table 3; all detected time is 200 s and 300 s for O<sub>2</sub>(0–1) and O(1S), respectively].

## 4. DETECTED NIGHTGLOW IER BY GBAIL

### A. Detected Nightglow IER Principle

The key components of the GBAIL prototype made by our group contain a wide-angle MI with a large air gap and a narrowband FPI [16]. The GBAIL is used to measure the upper atmospheric (80–120 km) wind velocity, temperature, VER, pressure, and gravitational wave. Some results are satisfactory [13,14]. The nightglow VER can also be inverted from the intensity of the imaging interference fringes of the GBAIL, which can be written as

$$I(\Delta) = I_0[1 + \exp(QT\Delta^2) \cos(2\pi\sigma_0\Delta)], \quad (16)$$

where *Q* = 1.82 × 10<sup>6</sup>/mλ<sub>0</sub><sup>2</sup>, λ<sub>0</sub> is wavelength at zero wind speed (λ<sub>0</sub> = 867.7 nm, or 557.7 nm), Δ is MI's total optical path difference (OPD), *I*<sub>0</sub> is the max intensity of the interference fringes, and *m* is the atomic mass. When the quantum numbers of the O<sub>2</sub> molecular vibrational-rotational spectrum vary, the max intensity *I*<sub>0</sub> in Eq. (16) is different, as it has been

modulated by the FPI of the GBAIL. The four-intensity-algorithm technique [6] requires the OPD Δ in Eq. (16) to be divided into two parts, i.e., fixed OPD Δ<sub>0</sub> and stepped OPD Δ' (Δ = Δ<sub>0</sub> + Δ' (Δ<sub>0</sub> ≫ Δ')). For the GBAIL, Δ<sub>0</sub> = 7.495 cm, and the stepped OPD Δ' = 0, λ<sub>0</sub>/4, 2λ<sub>0</sub>/4, and 3λ<sub>0</sub>/4 (λ<sub>0</sub> = 867.7 nm or 557.7 nm), respectively. Under the combined influence of the MI and FPI of the GBAIL, the group intensity for the quantum number *j* is

$$\begin{cases} I_{1(j)} = I_{0(j)}[1 + UV \cos \phi_{(j)}], \\ I_{2(j)} = I_{0(j)}[1 - UV \sin \phi_{(j)}], \\ I_{3(j)} = I_{0(j)}[1 - UV \cos \phi_{(j)}], \\ I_{4(j)} = I_{0(j)}[1 + UV \sin \phi_{(j)}], \end{cases} \quad (17)$$

where *U* and *V* are the resolutions of the GBAIL instrument and the interference fringe. With the four fringe intensities worked out, the max intensity of *I*<sub>0(j)</sub>, which is just related to the nightglow VER, can be obtained through Eq. (17) as

$$I_{0(j)} = [I_{1(j)} + I_{3(j)}]/2 = [I_{2(j)} + I_{4(j)}]/2, \quad (18)$$

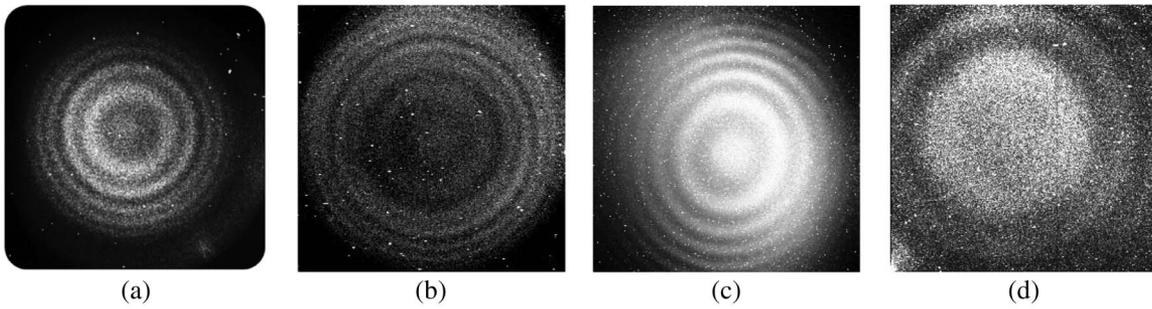
where the quantum number *j* corresponds to a rotational line. It should be pointed out that, for the single atom airglow of O(1S), the intensity *I*<sub>0</sub> only needs one fringe to invert the average IER, and the quantum number *j* is not a requisite. The intensities in Eq. (18) can be read by the GBAIL's CCD, whose imaging interference fringes are shown through the experiments in a later section. The operated intensity of Eq. (17) is just replaced by the average IER for one given fringe.

### B. Experiments by GBAIL

A series of urban and rural experiments were carried out with GBAIL in northwest China. Some of the GBAIL's pictures are selected to calculate VER. Four nightglow images of interference fringes for different dates and geographical positions are shown in Figs. 5(a)–5(d), where the white dots are stars or cosmic rays.

For the O<sub>2</sub>(0–1) 867.7 nm nightglow, the pictures of imaging interference fringes taken on 8 November 2011 and 6 December 2013 before daybreak at the urban area Xi'an in China, which is situated at N34°13'23" and E108°59'39", 457 m above sea level, are selected. One of the four stepped imaging interferometer fringes has been shot and is shown in Figs. 5(a) and 5(b). The CCD pixels are 512 × 512 (2 × 2 pixels for a bin), and the peak quantum efficiency is 0.38. The exposure time of each picture is 200 s.

Similarly, for the O<sub>2</sub>(0–1) 867.7 nm nightglow, the pictures of the imaging interference fringes shown in Fig. 5(c) were obtained on 18 April 2018 before daybreak at the rural area of the



**Fig. 5.** Imaging interference fringe of O<sub>2</sub>(0-1) 867.7 and O(<sup>1</sup>S) nightglow observed by the GBAIL (a) O<sub>2</sub>(0-1) at Xi'an University Qujiang campus (urban, N34°13'23", E 108°59'39", 8 November 2011; (b) O<sub>2</sub>(0-1) at Xi'an University Qujiang campus (urban), N34°13'23", E 108°59'39", 6 December 2013; (c) O<sub>2</sub>(0-1) at Ren zong miao (rural), N34°19'56", E109°16'53", 18 April 2018; (d) O(<sup>1</sup>S) at Ren zong miao (rural), N34°19'56", E109°16'53", 18 December 2011.

Zhuque Forest Park in China, which is located at N33°50'1" and E108°30'50", 1200 m above sea level. All the GBAIL's CCD shooting conditions and parameters in Fig. 5(c) are the same as those in Fig. 5(a). The exposure time of each picture is 200 s.

For the O(<sup>1</sup>S) 557.7 nm nightglow, the imaging interference fringes in Fig. 5(d) were obtained on 18 December 2011 before daybreak at the rural mountaintop of Ren zong Miao of Lin tong, which is situated at N34°19'56" and E109°16'53", 1302 m above sea level. All of the GBAIL's CCD shot conditions and parameters in Fig. 5(d) are the same as those in Fig. 5(a), except the exposure time of 300 s and altitude of 1302 m. The exposure time of each picture is 300 s.

**345 C. Retrieved Nightglow IER for GBAIL**

346 With the nightglow pictures shown in Figs. 5(a)–5(d) obtained,  
347 the following expression is used to calculate the detected  
348 nightglow IER<sub>D</sub> so as to derive the VER value:

$$IER_D = (N_s - N_d) \cdot \frac{1}{0.004^2} \cdot \frac{C_{ADV}}{\eta} \cdot \frac{1}{\xi} \cdot \frac{1}{\tau_{sys}} \cdot \frac{1}{t}, \quad (19)$$

349 where  $N_s$  is the electron counting read from one imaging  
350 interference fringe on the GBAIL's CCD, which should be  
351 the total count rate in counts,  $N_d$  is the dark current,  $t$  is ob-  
352 servation time [ $t$  is 200s for O<sub>2</sub>(0-1), 300 s for O(<sup>1</sup>S)],  $C_{ADV}$   
353 is the CCD's digital-to-analog conversion efficiency,  $\eta$  is the  
354 CCD's quantum efficiency,  $\xi$  is the P(0-1) branch pair of  
355 the lines' carrier value of the O<sub>2</sub>(0-1) band emission [22], and  
356  $\tau_{sys}$  is the GBAIL's transmittance. The value of 0.004 cm<sup>2</sup> is  
357 the united conversion of CCD pixel size (20 micron) to  
358 centimeters.

359 According to the factory parameters  $C_{ADV} = 2$  and  $\eta =$   
360 0.38 (for 867.7 nm wavelength) and  $\eta = 0.55$  (for 557.7 nm  
361 wavelength), the parameters of  $N_d$  and  $\tau_{sys}$  are detected in lab  
362 as  $N_d = 100$ ,  $\tau_{sys} = 0.18$ ,  $\xi = 7.5\%$ . By the software "Image  
363  $J$ ," one of the imaging interference fringes is used to read the  
364 electron count. Bringing these parameters and the count into  
365 Eq. (17), the nightglow IER of O<sub>2</sub>(0-1)867.7 nm can be ob-  
366 tained, and it is  $1.29 \times 10^7$  (photons · cm<sup>-2</sup> · s<sup>-1</sup>) (test date: 18  
367 April 2018). Other calculated and experimental results of the  
368 IER are shown in the fifth column of Table 3.

**D. Error and Analysis**

According to the calculated IER<sub>C</sub> and detected IER<sub>D</sub> in Eqs. (15) and (18), the relative errors are shown in the sixth column of Table 3.

Because our GBAIL is a ground-based device that measures the integral of the upper atmospheric VER profile along the line of sight, the number of photons emitted from a centimeter squared column along that line of sight every second. The minimum relative error of the VER of O<sub>2</sub>(0-1) measured by the GBAIL is 2.1%, the maximum error is 29.0%, and the error of the VER of O(<sup>1</sup>S) is 6.5% as shown in Table 3. Although the error for nitric oxide (NO) VER observed from the space-based SABER is 15% [12] and O<sub>2</sub> (0-0) VER error measured by HRDI is 10% [10], the GBAIL has accuracy limitations compared to other instruments. However, with mass data and suitable analysis this limitation may be reduced and the GBAIL's accuracy may be improved. We need more data to establish the daily variability in order to investigate the seasonal variability.

The relative errors are defined from comparisons with the simulations using NRLMSISE data. However, those are just differences with this model simulation, and it should be noted that they may have errors of every different nature (temperature, abundances, and pressures). Because the VERs are simulated and calculated based on the models of NRLMSISE-00 and MODTRAN in this paper, the model of "atmospheric visibility" only has the selected alternatives of 23 km or 10 km to reach the parameters in Eq. (5) and the MODTRAN model. However, our local detected data are in northwest China, which has dusty climate conditions. The geographical position in the present study is in northwest China where sandstorms occur more frequently, although the all experiments carried out by the GBAIL were selected to detect after rain. In fact, the actual atmospheric visibility cannot achieve the model's atmospheric visibility. Additionally, different aerosol scenarios should be considered, since the atmospheric transmittance is a factor of the column IER. We think this is one reason that may lead to the error between the calculated and detected IERs of the nightglows.

**5. CONCLUSIONS**

In order to obtain the upper atmospheric VER at 80–120 km altitude in northwest China, this paper studied the VER of

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O<sub>2</sub>(0-1) and O(<sup>1</sup>S) species nightglow based on the Barth mechanism, with the main conclusions being as follows.

The nightglow VER profile of O<sub>2</sub>(0-1) has been simulated and the peak values of VER at 94 km altitude on experimental dates 6 December 2013 and 8 November 2011 are 8111 and 8406 photons · cm<sup>-3</sup> · s<sup>-1</sup>, respectively. The nightglow emission will be attenuated by the atmosphere, and its simulated transmittances of O<sub>2</sub>(0-1) for 8 November 2011 are 0.670 (0.671) and 0.567 (0.560) when the zenith angles are 0° and 45°, respectively. After the nightglow VER is changed into the IER, the theoretical calculation nightglow IER<sub>C</sub> values of O<sub>2</sub>(0-1) for 0° and 45° zenith angles, respectively, are 1.48 × 10<sup>7</sup> and 1.91 × 10<sup>7</sup> photons · cm<sup>-2</sup> · s<sup>-1</sup>. The detected IER<sub>D</sub> values of O<sub>2</sub>(0-1) nightglow for 45° zenith angle by the GBAIL on 8 November 2011 and 6 December 2013 are 1.87 × 10<sup>7</sup> and 1.27 × 10<sup>7</sup> photons · cm<sup>-2</sup> · s<sup>-1</sup>, respectively.

The nightglow VER profile of O(<sup>1</sup>S) has been simulated and the peak value of VER is 338 photons · cm<sup>-2</sup> · s<sup>-1</sup> at 96 km. The nightglow emission is going to be attenuated by the atmosphere, and the simulated transmittances of O(<sup>1</sup>S) for 0° and 45° zenith angles, respectively, are 0.658 and 0.551 on 18 December 2011. After the nightglow VER is changed into the IER, the theoretical nightglow IER<sub>C</sub> values of O(<sup>1</sup>S), respectively, are 5.53 × 10<sup>5</sup> and 7.03 × 10<sup>5</sup> photons · cm<sup>-2</sup> · s<sup>-1</sup> for 0° and 45° zenith angles. The IER<sub>D</sub> value of O(<sup>1</sup>S) nightglow detected by GBAIL is 6.57 × 10<sup>5</sup> photons · cm<sup>-2</sup> · s<sup>-1</sup>.

**6 Funding.** National Natural Science Foundation of China (NSFC) (61675165); Natural Science Foundation of Shaanxi Province (2016JM1011); National Instrument Development Special Program (2013YQ030651); Characteristic Fund of Xi'an University of Technology (2015TS012); Open Research Fund of Key Laboratory of Spectral Imaging Technology, Chinese Academy of Sciences (LSIT201714D).

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# Queries

1. AU: Edit to title OK?
2. AU: Please check if there is missing text at the end of this sentence after “from”: “where the unit of IERC is photons•cm<sup>-2</sup>•s<sup>-1</sup>, the unit of *h* is denoted by kilometers (km), and the factor of 105 is the united unit result from.”
3. AU: Please check the edit made to the following sentence: “ $\xi$  is the P(0-1) branch pair of the lines carrier value of the O2(0-1) band emission [22].”
4. AU: Please check if there is missing text at the end of this sentence: “Because our GBAll is a ground-based device that measures the integral of the upper atmospheric VER profile along the line of sight, the number of photons emitted from a centimeter squared column along that line of sight every second.”  
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5. AU: Please define “SABER” in this sentence: “Although the error for nitric oxide (NO) VER observed from the space-based SABER is 15%”
6. AU: The funding information for this article has been generated using the information you provided to OSA at the time of article submission. Please check it carefully. If any information needs to be corrected or added, please provide the full name of the funding organization/institution as provided in the CrossRef Open Funder Registry (<https://search.crossref.org/funding>).  
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